

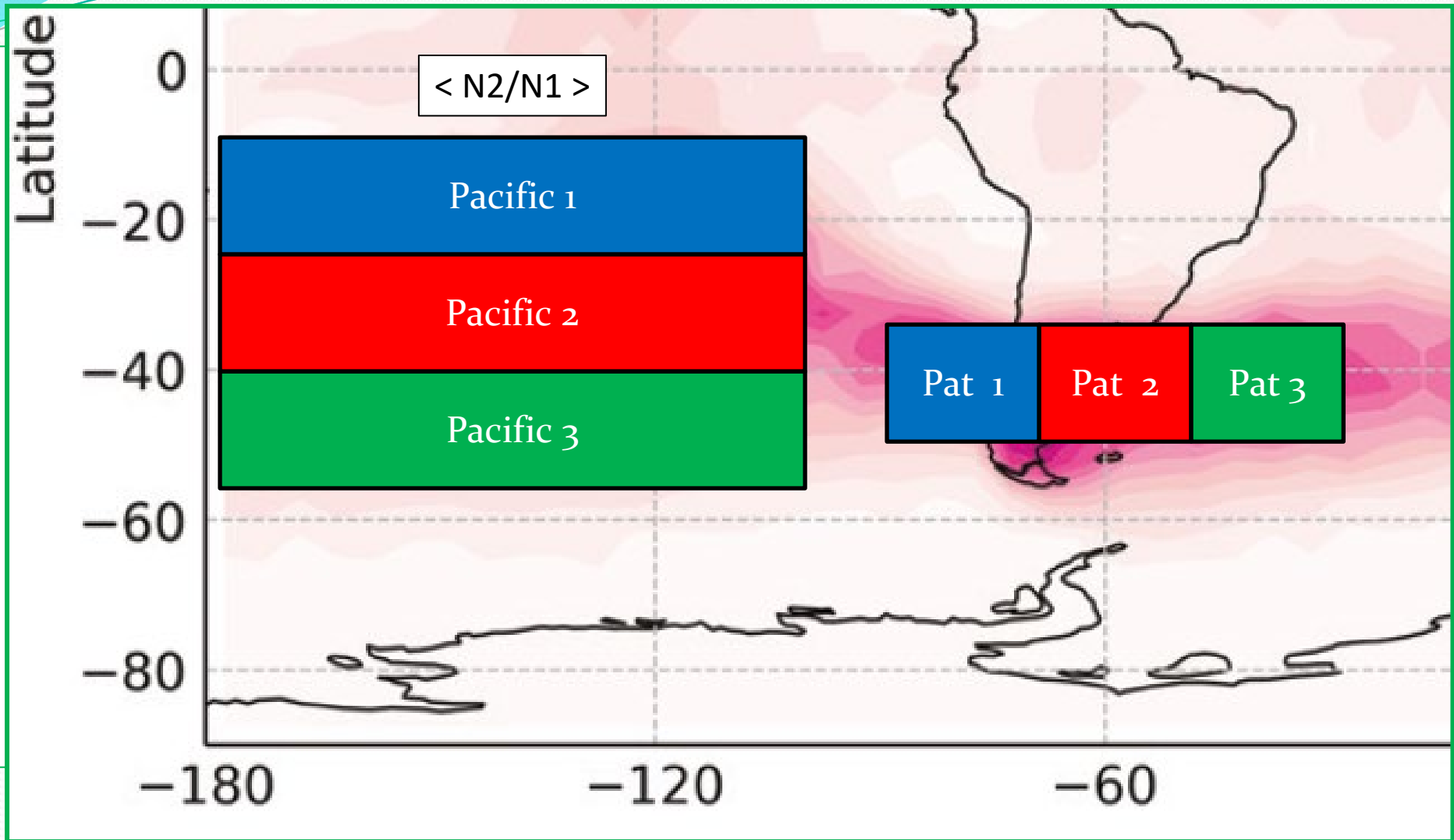
# Signs of climate variability in double tropopause global distribution from two decades of radio occultation data

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# Some previous knowledge

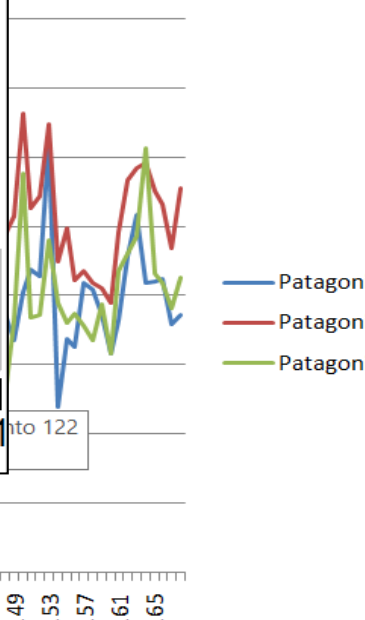
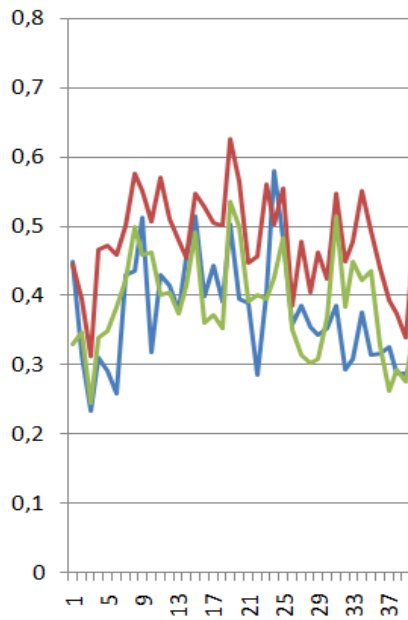
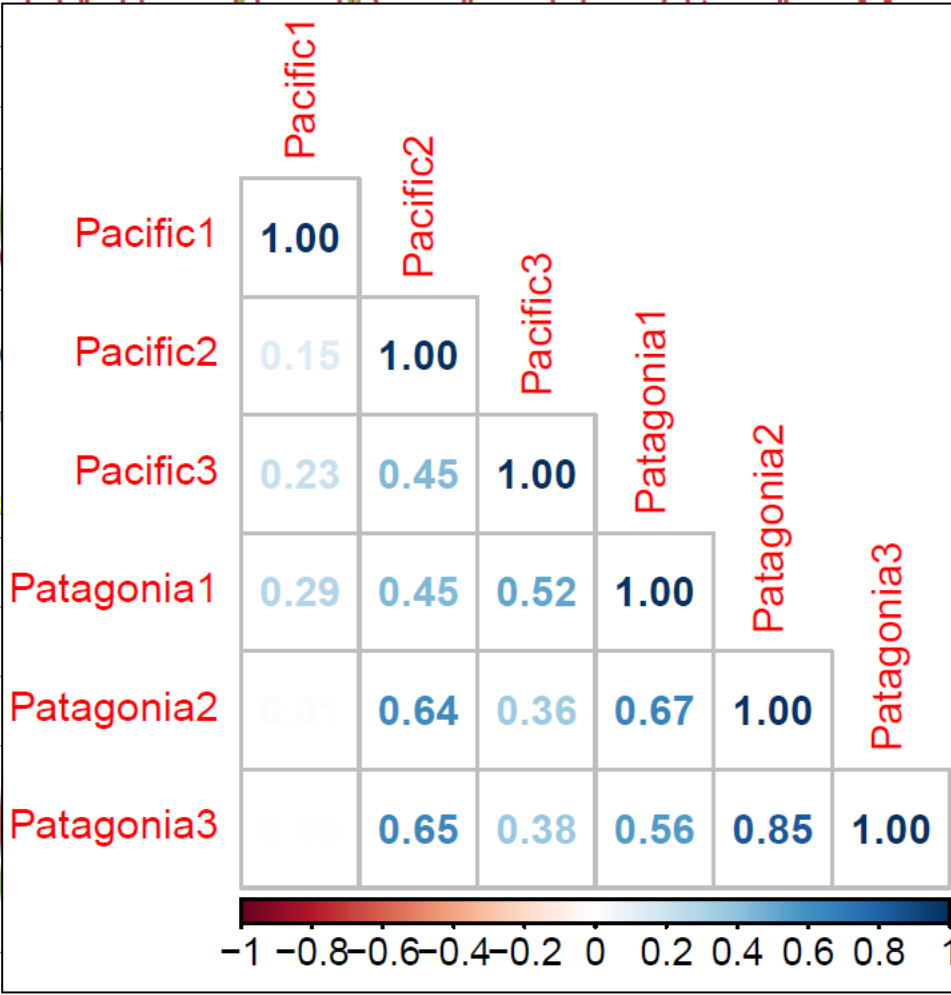
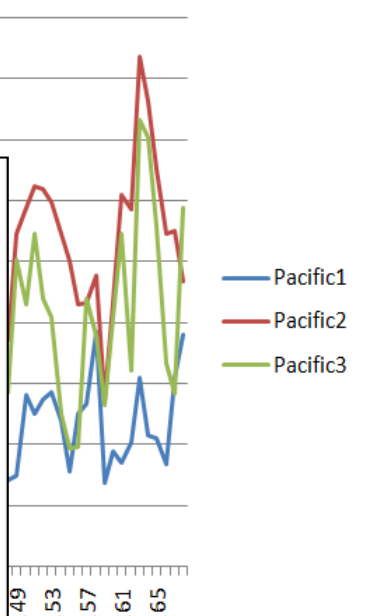
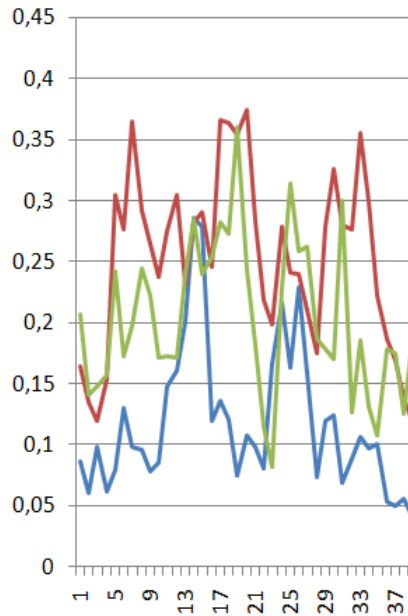
- vertical folding of the tropopause, occurrence of DTs
- midlatitudes (Peevey et al., 2014; Wang & Polvani, 2011, Homeyer et al., 2010; Castanheira et al., 2012; Liu & Barnes, 2018; Pan et al., 2009; Randel et al., 2007)
- storm track regions, subtropical jet stream (Bischoff et al., 2007; Schmidt et al., 2006; Seidel & Randel, 2006)
- mountain gravity waves (Schmidt et al., 2006) and
- cyclogenesis (Añel et al., 2008)
- Brewer-Dobson circulation (Castanheira et al., 2012)
- cloud-top inversion layers (Biondi et al., 2012)
- radiosonde measurements
- RO data (e.g. Randel et al., 2007; Schmidt et al., 2006; Xu et al., 2014) including the relationship to ENSO events (Wilhelmsen et al JGR 2020)



$$\text{DTs percentage} = (N2/N1) \times 100$$

(from Wilhelmsen et al, 2020)

< N2/N1 >



month number

# A preliminary work:

As an approximation to relate a meteorological signal using different indices,

Llamedo et al., 2016

(after Randel & Wu, 2015 and other relevant contributions)

ENSO, PW, q and T

$$\begin{aligned} F(t) = & A_0 + A_1 t + A_2 \sin(\omega t) + A_3 \cos(\omega t) + A_4 \sin(2\omega t) \\ & + A_5 \sin(2\omega t) + A_6 \sin(3\omega t) \\ & + A_7 \cos(3\omega t) + B_1 \text{QBO1}(t) \\ & + B_2 \text{QBO2}(t) + B_3 \text{SF}(t) \end{aligned}$$

# Multiple linear regression

A multiple regression model is written as

$$\begin{aligned}y_i &= \beta_0 + \sum_{j=1}^d x_{ij}\beta_j + \epsilon_i \\ &= \beta_0 + \beta_1x_{1i} + \beta_2x_{2i} + \dots + \beta_dx_{di} + \epsilon_i, \quad i = 1, 2, \dots, n\end{aligned}$$

Climate indices from :  
[psl.noaa.gov/data/climateindices/list](https://psl.noaa.gov/data/climateindices/list)

The  $\beta$  coefficients are obtained by minimizing the residual sum of squares, thus obtaining:

$$\hat{\beta} = (\mathbf{X}^T \mathbf{X})^{-1} \mathbf{X}^T \mathbf{y}$$

Given new data  $\mathbf{X}_{new}$ , the least squares prediction is:

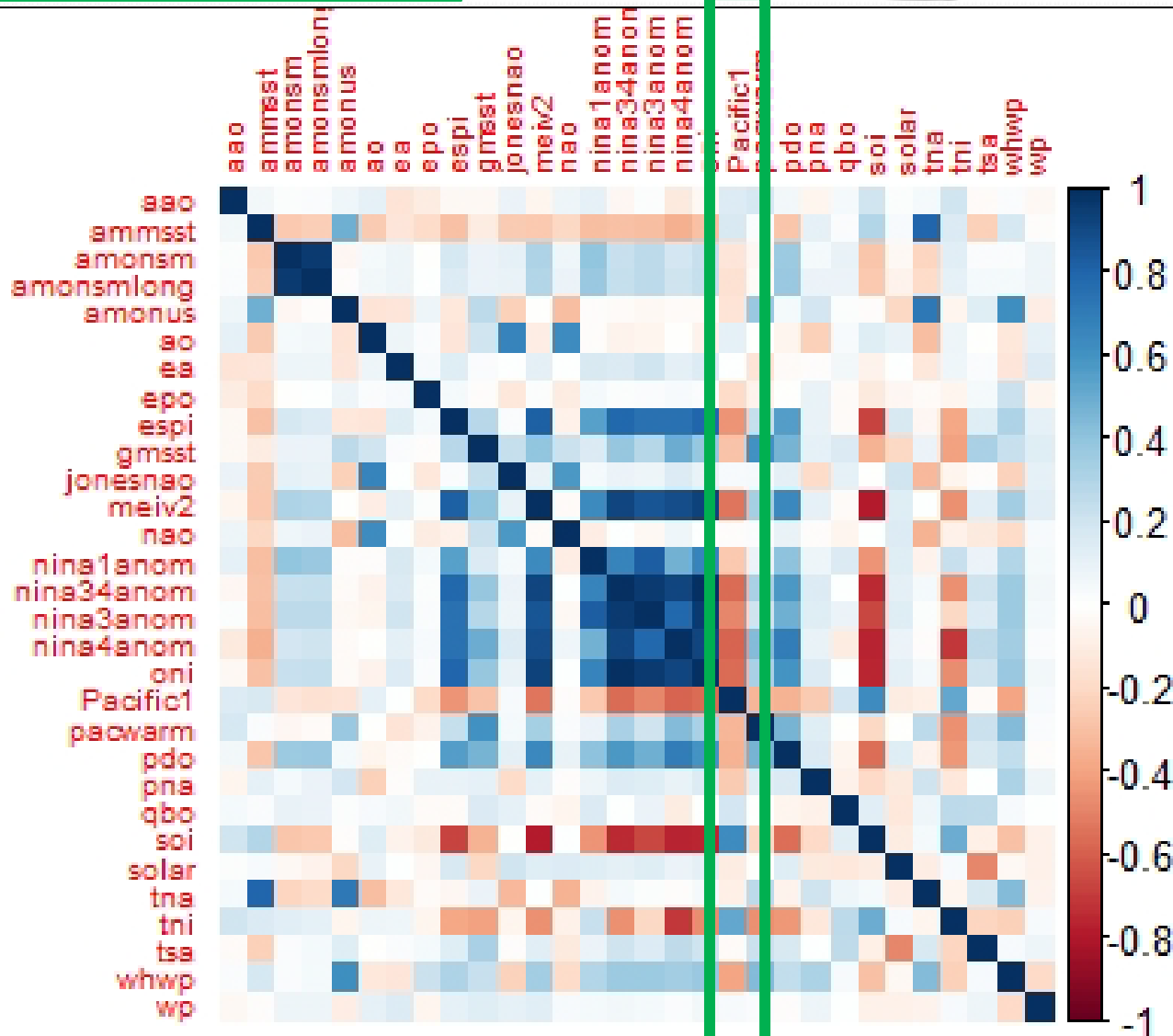
$$\hat{y} = \mathbf{X}_{new} \hat{\beta} = \mathbf{X}_{new} (\mathbf{X}^T \mathbf{X})^{-1} \mathbf{X}^T \mathbf{y}$$

$$RMSE = \sqrt{\frac{\sum_{i=1}^n (pred_i - obs_i)^2}{n}}$$

Root Mean Squared Error

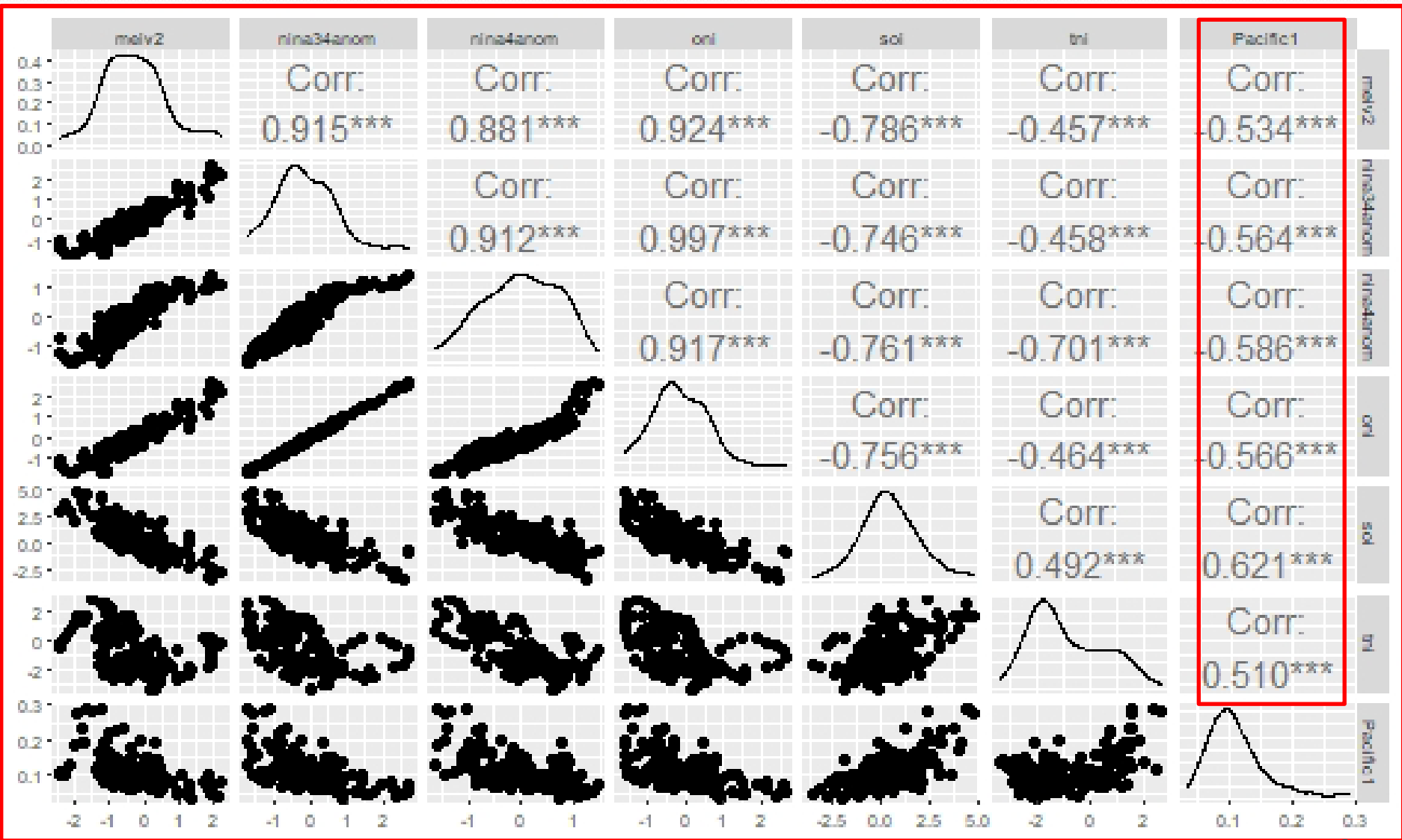
e.g.: Pacific 1 subregion

Climate indices from :  
[psl.noaa.gov/data/climateindices/list](http://psl.noaa.gov/data/climateindices/list)



Correlations with  $R^2 > 0.5$  is found between the response variable and predictors:  
"meiv2" , "nina34anom" , "nina4anom" , "oni" , "soi" , "tni"

Considering only these latter “best” predictors  
 (“meiv2”, “nina34anom”, “nina4anom”, “oni”, “soi”, “tni”) for “Pacific1”:



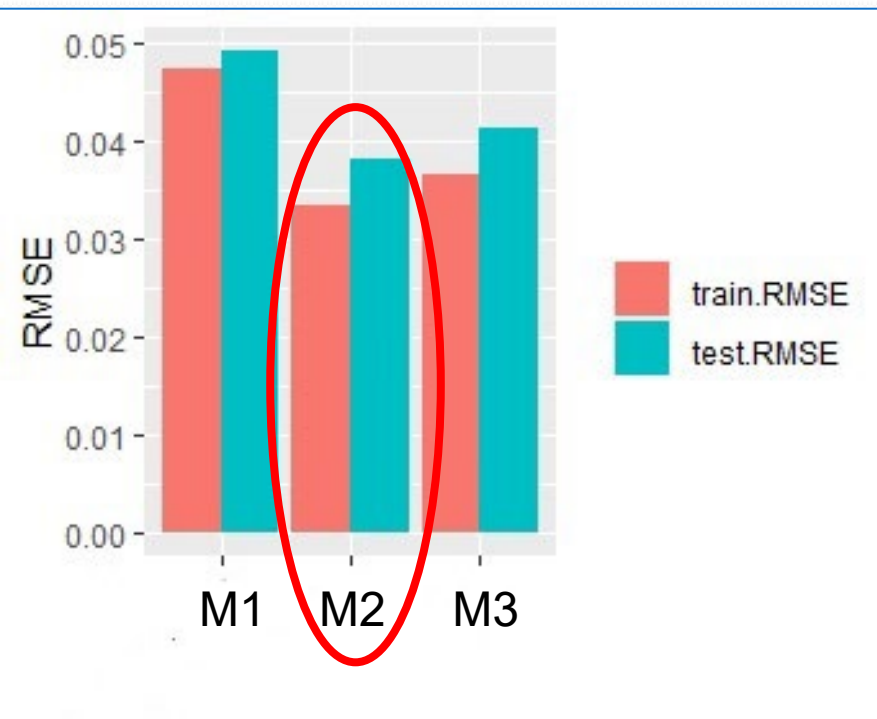


Three regression models to predict DTs from different sets of arbitrary predictors were compared:

**M1)** “baseline” (only  $\beta_0$ )

**M2)**  $\beta_0 + \beta_1 \text{soi} + \beta_2 \text{whwp} + \beta_3 \text{tni} + \beta_4 \text{oni}$

**M3)** “linear” ( $\beta_0 + \beta_1 \text{soi}$ )



- As an example, for **M2**:

- Residuals:

Min	1Q	Median	3Q	Max
-0.0677	-0.0234	-0.0030	0.0171	0.1009

- Coefficients:

	Estimate	Std. Error	t value
(Intercept)	0.118982	0.004959	23.993
soi	0.011925	0.003033	3.931
whwp	-0.004230	0.001483	-2.852
oni	-0.005017	0.005591	-0.897
tni	0.007947	0.002397	3.316

model	train RMSE	test RMSE
M1	0.04759141	0.04926829
→ M2	0.03345134	0.03828805
M3	0.03656820	0.04127345

# Summary

- 5 sub-regions (Pacific 2-3 and Patagonia 1-2-3) show a correlation between their respective monthly averaged TDs
- DTs in *Pacific 1* (to the north of the subtropical jet) exhibit a different behavior
- However, it correlates better with the global indices than the other 5 subregions do
- Progressive addition of global indices to the model improves its performance, reducing the RMSE of training and testing.

# Current work

- A possible correlation between the observed distribution of DTs (from RO) and inertio gravity waves, usually radiated by geostrophic adjustment near to the subtropical jet (predictor proposed: cross-current Rossby number) is being analyzed.
- A principal component study is being performed on the available 20-year DTs dataset and 5x5 lat/lon deg pixels in the southern hemisphere.

*Thank you for your attention*

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