Accepted version of the article:

Reimann, T., Rehrl, C., Shoemaker, W. B., Geyer, T., Birk, S. (2011): The significance of turbulent flow representation in single-continuum models. Water Resources Research 47, W09503, doi:10.1029/2010WR010133.

To view the published open abstract, go to <u>http://dx.doi.org</u> and enter the DOI.

An edited version of this paper was published by AGU.

Copyright (2011) American Geophysical Union.

1	The significance of turbulent flow representation in single-continuum
2	models
3	Thomas Reimann <sup>1,*</sup> , Christoph Rehrl <sup>2</sup> (†), W. Barclay Shoemaker <sup>3</sup> ,
4	Tobias Geyer <sup>4</sup> , Steffen Birk <sup>2</sup>
5	1,*) Corresponding author: Thomas Reimann, TU Dresden, Institute for Groundwater
6	Management, D-01062 Dresden, Germany, email: Thomas.Reimann@tu-dresden.de, phone:
7	+49 351 46342555, Fax: +49 351 46342552;
8	2) University of Graz, Institute for Earth Sciences, Heinrichstrasse 26, A-8010 Graz, Austria,
9	email: steffen.birk@uni-graz.at, phone: +43 316 380 5583;
10	3) U.S. Geological Survey, 3110 SW 9th Ave, Fort Lauderdale, Florida 33315, USA, email:
11	bshoemak@usgs.gov, phone: +1 954 377 5956;
12	4) University of Göttingen, Geoscientific Centre, Goldschmidtstrasse 3, D-37077 Göttingen,
13	Germany, email: tgeyer@gwdg.de, phone: +49 551 39 9398
14	Keywords: karst hydrology; groundwater flow; numerical modeling; dual porosity; turbulent
15	flow

16 AGU Index Terms: 1828 Groundwater hydraulics

#### 17 Abstract

18 Karst aquifers exhibit highly conductive features caused from rock dissolution processes. 19 Flow within these structures can become turbulent and therefore be expressed by nonlinear 20 gradient functions. One way to account for these effects is by coupling a continuum model 21 with a conduit network. Alternatively, turbulent flow can be considered by adapting the 22 hydraulic conductivity within the continuum model. Consequently, the significance of 23 turbulent flow on the dynamic behavior of karst springs is investigated by an enhanced single-24 continuum model that results in conduit-type flow in continuum cells – CTFc. The single-25 continuum approach CTFc represents laminar and turbulent flow as well as more complex 26 hybrid models that require additional programming and numerical efforts. A parameter study 27 is conducted to investigate effects of turbulent flow on the response of karst springs to 28 recharge events using the new CTFc approach, existing hybrid models, and 29 MODFLOW-2005. Results reflect the importance of representing (1) turbulent flow in karst conduits and (2) the exchange between conduits and continuum cells. More specifically, 30 31 laminar models overestimate maximum spring discharge and underestimate hydraulic 32 gradients within the conduit. It follows that aquifer properties inferred from spring 33 hydrographs are potentially impaired by ignoring flow effects due to turbulence. The 34 exchange factor used for hybrid models is necessary to account for the scale dependency 35 between hydraulic properties of the matrix continuum and conduits. This functionality, which is not included in CTFc, can be mimicked by appropriate use of the Horizontal Flow Barrier 36 37 package for MODFLOW.

## 38 **1. Introduction**

39 Karst aquifers can be conceptualized as dual-flow systems comprised of (1) fractured porous rock-matrix (defined herein as matrix) with high storage-capacity, which includes 40 41 primary or intergranular porosity and secondary porosity due to fractures, and (2) highly permeable solution conduits (tertiary porosity) with relatively high-velocity flow and low 42 43 storage-capacity. These conduits form together with the surrounding matrix a complex and 44 heterogeneous flow system. Commonly, flow in the low permeability matrix is laminar. 45 Conversely, flow in the discrete and localized solution conduits is rapid and therefore often 46 turbulent. Hydraulic signals, such as pulse discharge due to recharge events, are rapidly 47 transmitted through solution conduits both in pressurized and open-channel flow [e.g. 48 Covington et al., 2009]. Karst springs often respond quickly and strongly to recharge events. 49 During recession periods, however, conduit flow is sustained by draining the storage provided 50 by the matrix [Ford and Williams, 2007]. To simulate responses of hydraulic heads and spring 51 discharge to recharge events, the solution conduit and matrix flow processes need to be linked. Numerical models that incorporate this linkage are often used for this purpose. 52 53 Appropriate modeling approaches frequently employed for this reason are described by 54 Teutsch and Sauter [1998].

55 In accordance with the dual-flow conceptual model, flow in the matrix can be treated 56 as a continuum flow field, whereas flow in the conduit system is simulated using a discrete 57 pipe network that accounts for both laminar and turbulent flow conditions. Hybrid models 58 (HM) coupling continuum and discrete approaches are frequently employed for generic 59 modeling in basic research, e.g., to simulate and analyze flow and transport processes [for example Király, 1984; Eisenlohr et al., 1997; Birk et al., 2006] or the mechanism of 60 61 speleogenesis [for example Liedl et al., 2003; Rehrl et al., 2008; Kaufmann, 2009]. Recently, Hill et al. [2010] applied the United States Geological Survey (USGS) hybrid model Conduit 62

Flow Process Mode 1 (CFPM1) [Shoemaker et al., 2008a] to simulate transient flow in the karst aquifer of Weeki Wachee, in west central Florida. In general, hybrid models are rarely applied to real aquifer systems as they require detailed information on aquifer geometry, hydraulic parameters, and boundary conditions, which often are not available. Further, the computational efforts of such hybrid models are considerable and the practical compatibility with existing model-related tools like graphical user interfaces, calibration tools, or other model packages is fairly limited.

70 On the contrary, single-continuum models (SCM) are more frequently used for practical applications because they offer several advantages, e.g. lower parameter 71 requirements, reduced numerical demands, and easily accessible codes. Such models 72 73 represent conduit influenced areas by highly conductive cells (smeared conduit approach; 74 Painter et al. [2006], Worthington [2009]) within an SCM. Commonly, groundwater flow in 75 SCMs such as MODFLOW-2005 [Harbaugh, 2005] is computed using the Darcy equation, which considers laminar flow only and therefore ignores potential effects of turbulence. 76 Scanlon et al. [2003] used an SCM to simulate regional groundwater flow at Barton Springs 77 Edwards aquifer (Texas). The authors conclude that the model cannot, however, accurately 78 79 simulate local directions or rate of groundwater flow because major conduits are not explicitly 80 designated and turbulent flow is not represented. Lindgren et al. [2004] also simulated the 81 Edwards aquifer (Texas) with a laminar SCM and found similar results, concluding that the 82 incorporation of turbulent flow could facilitate a better simulation of groundwater flow and 83 transport.

One alternative approach to apply SCMs on a local scale is to represent conduits with relatively small model cells that allow laminar and turbulent flow. The Conduit Flow Process Mode 2 (CFPM2) [Shoemaker et al., 2008a] for MODFLOW-2005 [Harbaugh, 2005] is a promising new method that accounts for turbulent flow in the continuum. Originally, this

approach was intended to reflect flow conditions in highly permeable, stratiform porous layers ("vuggy layers") [Shoemaker et al., 2008a, b; Kuniansky et al., 2008]. A newly developed modification of CFPM2 considers turbulent flow more generally using a user defined nonlinearity of flow [Reimann et al., 2011]. Setting the parameter that controls the nonlinearity to an appropriate value results in conduit-type flow in continuum cells, subsequently shortened as CTFc.

94 The objective of this work is to answer the following questions: (1) Is it possible to 95 simulate groundwater flow affected by strong local heterogeneities with a continuum 96 approach that accounts for laminar and turbulent flow? (2) Is this laminar and turbulent 97 continuum model appropriate to reflect the dual-flow behavior of karst aquifers, resulting 98 from matrix-conduit interaction? (3) What is the relevance of turbulent conduit flow with 99 respect to the dynamic responses typically observed at karst springs? To consider these 100 questions, SCM approaches are compared with HM approaches that are known to provide 101 adequate representations of the dual-flow behavior of karst aquifers [Birk et al., 2005; Hill et 102 al., 2010]. The SCM approaches employed include the aforementioned CFPM2, which is 103 intended to represent turbulent flow in vuggy layers, and the modified version of CFPM2 104 representing CTFc.

105 To this end, equations to describe turbulent flow in discrete structures are assembled. 106 Subsequently, the flow equation originally intended for the laminar continuum is re-derived to analogously consider turbulent conduit-type flow - CTFc - with appropriate parameters. 107 108 Next, CTFc is applied to simulate a coupled matrix-conduit system representing a 109 hypothetical karst catchment. In extension to Covington et al. [2009], who examined the 110 transmission of recharge pulses through single elements of karst aquifers, this application 111 considers the signal transmission through a coupled conduit-matrix system. CTFc results are 112 validated by comparison with existing numerical approaches that explicitly consider turbulent

113 conduit flow. Furthermore, the influence of the matrix as well as the conduit-matrix 114 interaction on conduit flow was studied. Finally, the article discusses the differences between 115 the turbulent continuum approach CTFc and existing continuum (MODFLOW-2005) as well 116 as hybrid (CFPM1) models with emphasis on future applications.

## 117 **2. Description of Karst Aquifer Hydraulics**

118 Governing laminar and turbulent flow equations are presented with the aim of 119 incorporating turbulent flow in continuum cells. The transition between laminar and turbulent 120 flow is governed by the dimensionless Reynolds number [Bear, 1988]:

121

122

$$\operatorname{Re} = \frac{\operatorname{qd}^*}{\operatorname{\nu}} \tag{1}$$

123

where q denotes the specific discharge [L/T] defined as discharge per unit cross section flow area, d\* is some specific length dimension [L] (mean void diameter of porous media or conduit diameter for discrete elements), and v is the kinematic viscosity of the fluid [L<sup>2</sup>/T].

127 2.1 Flow Processes in the Matrix

128 Slow and laminar groundwater flow in the matrix along the i-th direction is computed129 using the linear Darcy equation,

130

131 
$$q_i = -K_i \frac{\partial h}{\partial x_i}$$
(2)

132

where K is the hydraulic conductivity [L/T], h represents the hydraulic head [L], and x is the spatial coordinate along flow direction [L]. The Darcy equation is valid for laminar flow as long as the Reynolds number does not exceed a value of around 1 to 10 [Bear, 1988]. 136 More generally, flow processes can be described with polynomial laws. A universal137 formulation is presented by Muskat [1946]

138

139 
$$\frac{\partial \mathbf{h}}{\partial \mathbf{x}_{i}} = \beta_{1} \mathbf{q}_{i} + \beta_{2} \mathbf{q}_{i}^{m}$$
(3)

140

141 where m represents the constant exponential term of the power law, which is assumed to take values between 1 and 2 [Sen, 1995], and  $\beta_1$  [T/L] as well as  $\beta_2$  [T<sup>m</sup>/L<sup>m</sup>] are parameters related 142 to fluid and rock properties. The linear term of the polynomial covers laminar flow and can be 143 144 interpreted as the Darcy equation with hydraulic conductivity  $K = 1/\beta_1$ . The nonlinear term describes turbulent flow by a power law. Several empirical and semi-empirical power laws 145 146 can be derived from this equation to account for turbulent flow [Sen, 1995]. A common 147 expression was suggested by Forchheimer [1901] with m = 2.0. For turbulent flow conditions 148 (Re >> 10) with high specific discharge the linear term in equation 3 becomes small and 149 therefore can be neglected.

# 150 2.2. Flow Processes in Conduits

Subsequently, equations for discrete flow in cylindrical conduits with circular crosssection are described in order to transfer the characteristics to the continuum flow model. Existing HMs like CFPM1 [Shoemaker et al., 2008a] or CAVE [Liedl et al., 2003] consider one-dimensional flow processes in karst conduits using the Darcy-Weisbach approach [Young et al., 2004]

156

157 
$$\frac{\partial h}{\partial x} = -\mu \frac{q^2}{2gd}$$
(4)

where  $\mu$  represents the friction coefficient [-], which controls the flow regime, d is the conduit diameter, and g denotes the gravitational acceleration [L/T<sup>2</sup>]. Laminar flow in cylindrical conduits is described with the linear Hagen-Poiseuille equation that can be obtained from the Darcy-Weisbach approach if  $\mu = 64/\text{Re}$  [Young et al., 2004]

163

164 
$$q = -\frac{gd^2}{32\nu} \frac{\partial h}{\partial x}$$
(5)

165

166 Therefore, laminar flow in the matrix and the conduits (compare equation 2) is described with167 structurally similar equations resulting in equivalent flow characteristics (Figure 1).

A critical Reynolds number  $\text{Re}_{c}$  of about 2,000 [Jain, 2001] characterizes the transition from laminar ( $\text{Re} < \text{Re}_{C}$ ) to turbulent flow conditions ( $\text{Re} > \text{Re}_{C}$ ) in cylindrical pipes. Furthermore, turbulent flow conditions can be subdivided into a transition zone (slightly turbulent with Reynolds numbers moderately exceeding  $\text{Re}_{C}$ ) and a fully turbulent zone [Jain, 2001]. Turbulent flow in cylindrical conduits can be represented by the Colebrook-White equation [Young et al., 2004], which is also based on the Darcy-Weisbach approach (equation 4)

175

176 
$$q = -2\log\left(\frac{k_c}{3.71d} + \frac{2.51\nu}{d\sqrt{2gd\frac{\partial h}{\partial x}}}\right)\sqrt{2gd\frac{\partial h}{\partial x}}$$
(6)

177

178 where  $k_c$  denotes roughness [L]. The Colebrook-White equation covers flow in the transition 179 zone as well as in the fully turbulent zone. The relation between specific discharge q and 180 hydraulic gradient  $(\partial h / \partial x)$  is

182 
$$q \propto \log \left( c + \frac{1}{\sqrt{\frac{\partial h}{\partial x}}} \right) \sqrt{\frac{\partial h}{\partial x}}$$
 (7)

with c representing some constant value. The logarithmic term in equation 7 approaches a constant value (log c) with increasing hydraulic gradient resulting in similarity to the nonlinear term of the generalized power law with m = 2.0 (equation 3).

187 Alternatively, turbulent flow can be described with Manning's equation [Jain, 2001],
188 which is primarily intended for free-surface flow but is also applicable for karst conduits, for
189 example Peterson and Wicks [2006] and Meyer et al., [2008]

190

191 
$$q = \frac{1}{n} \left( r_{hy} \right)^{\frac{2}{3}} \sqrt{\frac{\partial h}{\partial x}}$$
(8)

192

where n represents the Manning coefficient  $[T/L^{(1/3)}]$  and  $r_{hy}$  is the hydraulic radius [L] (equal to cross section flow area divided by wetted perimeter). The Manning equation considers turbulent flow in the fully turbulent zone by a power law relation

196

197 
$$q \propto \sqrt{\frac{\partial h}{\partial x}}$$
 (9)

198

and, therefore, equals the nonlinear term of the generalized power law (equation 3) with an exponential term m = 2.0. Because Manning's equation is widespread in environmental hydraulics, several references regarding parameterization are available (Table 1). 202 **Table 1:** Values for Manning coefficient n.

Description	Value for Manning coefficient n [s/m <sup>(1/3)</sup> ]	Author
Natural channel, earth, smooth	0.020	Jain [2001]
Natural channel with gravel beds, straight	0.025	Jain [2001]
Natural channel with gravel beds, with boulders	0.040	Jain [2001]
Karst conduits (Devil's Icebox – Connor's Cave System in central Missouri, model application)	0.035	Peterson and Wicks [2006]

Figure 1 illustrates the relation between specific discharge and hydraulic gradient for an example situation (d = 0.2 m,  $k_c = 0.01$  m, n = 1.83 x 10<sup>-2</sup> s/m<sup>(1/3)</sup>). For considerably turbulent conditions the Colebrook-White equation equals the Manning equation and also the generalized power law with m = 2.0. Consequently, the Colebrook-White and Manning equation can be converted into each other using the relation [Jain, 2001]

209

$$n \propto k_c^{\frac{1}{6}}$$
 (10)

211

210

where the adequate conversion can be found by empirical investigations, for instance as  $n = (1/24) k_c^{(1/6)}$  [Jain, 2001]. Hence, for hydraulic situations typically found in karst systems the Colebrook-White and the Manning equation produce similar results (compare equations 7, 9, and Figure 1).



Figure 1: Hydraulic gradient and corresponding discharge for different linear (laminar flow) and nonlinear (turbulent flow) equations. The example is computed for a conduit with d = 0.2m,  $k_c = 0.01$  m, and  $n = 1.83 \times 10^{-2}$  s/m<sup>(1/3)</sup>.

While transitioning from laminar to turbulent flow and vice versa, conduit flow may be 220 221 influenced by hysteresis effects, as laminar flow tends to stay laminar and turbulent flow 222 tends to stay turbulent. This effect is observed for pipes by several authors [Kanda and 223 Yanagiya, 2008]. The hysteresis process is described by two critical Reynolds numbers as 224 shown in Figure 1. Re<sub>C1</sub> denotes the transition from laminar to turbulent flow conditions and 225  $Re_{C2}$  represents the transition from turbulent to laminar flow conditions. Depending on the 226 values for Re<sub>C1</sub> and Re<sub>C2</sub>, discharge may exhibit a step-wise behavior while transitioning 227 between these two regimes. Hence, these two critical Reynolds numbers are considered 228 during laminar and turbulent flow computation.

#### 229 2.3. Interaction between Matrix and Conduit Flow

230 Both matrix and conduits interact with each other resulting in the aforementioned dual 231 flow characteristics. The hydraulic parameters of the matrix are influenced by small scale 232 fissures and fractures. The hydraulic conductivity of the matrix at a field-relevant length scale, therefore, is expected to be higher than that at laboratory scale (0 - 1 m) [Kiraly, 2002]. This 233 234 scale-dependency of hydraulic properties needs to be addressed when investigating the 235 interaction between matrix and karst conduits. Figure 2 shows the cross-section of a 236 conceptual model of a karst conduit draining the matrix and the associated flow components 237 and capacities. The different flow components are described in more detail below. This 238 conceptual interpretation is based on several assumptions as subsequently denoted and aims to 239 give a simplified overview of the importance of the individual components.



Figure 2: Cross section of a karst conduit embedded in the matrix. The flow system is controlled by matrix flow to karst conduits  $(Q_m)$ , conduit infiltration capacity  $(Q_i)$ , and conduit flow capacity  $(Q_c)$ .

 $(Q_m/L)$  Flow in the matrix toward karst conduits and vice versa is the volumetric flow rate provided by the matrix due to the hydraulic gradient in the matrix induced by the conduit. This flow can be approximated with the equations for steady flow in an unconfined homogeneous aquifer as [Bear, 1988]

248

249 
$$\frac{Q_m}{L} = K_{m1} \frac{h_0^2 - h_1^2}{\Lambda}$$
(11)

250

where  $Q_m/L$  represents matrix flow to or from the karst conduit  $[L^2/T]$  per unit length,  $K_{m1}$  is the hydraulic conductivity of the matrix for a sufficient length scale [L/T],  $h_0$  is the hydraulic head [L] at the distance  $\Lambda$  [L] perpendicular to the conduit flow direction [L] and  $h_1$  is the hydraulic head [L] in the matrix at the conduit. Note that equation 11 covers the flow to the conduit from both sides.

256  $(Q_i/L)$  The conduit infiltration capacity is the volumetric flow rate per unit length that can be 257 gathered by the cylindrical karst conduit based on Darcy's equation

258

$$\frac{Q_i}{L} = \pi dK_{m2} I_{m2} \tag{12}$$

260

where  $Q_i/L$  represents the infiltration capacity of the conduit  $[L^2/T]$  per unit length,  $K_{m2}$  is the hydraulic conductivity of the matrix in the direct vicinity of the conduit [L/T], and  $I_{m2}$  denotes the dimensionless inflow gradient between karst conduit and matrix. The conduit infiltration capacity can be affected by clastic, organic and precipitated sediments deposited in a karst conduit, e.g. wall coatings and crystal growth (details about cave interior deposits can be found in Ford and Williams [2007]). In general,  $K_{m2}$  is smaller than  $K_{m1}$  due to increased influence of fissures and fractures at larger scale as mentioned above.  $\frac{(Q_c) \text{ Conduit flow capacity}}{(Q_c) \text{ Conduit flow capacity}} \text{ is calculated for laminar flow conditions with a linear flow}$ equation (e.g., Hagen-Poiseuille equation) and for turbulent flow conditions with nonlinearflow equations (e.g., Colebrook-White equation, Manning equation). Consequently, thehydraulic gradient inside the conduit I<sub>c</sub> depends on discharge, which represents the balance ofinflow and outflow per unit length along the conduit flow direction. Accordingly, the conduitflow capacity affects the infiltration capacity Q<sub>i</sub>/L.

274 The flow components Q<sub>m</sub>/L, Q<sub>i</sub>/L, and Q<sub>c</sub> control the event-induced hydraulic signal 275 transmission in terms of discharge and head in both the karst conduit and the matrix. 276 Depending on the amount of water gathered by the conduit, a hydraulic gradient in the matrix (approximated as I<sub>m1</sub>), which drives water toward the draining conduit, and a hydraulic 277 278 gradient between conduit and matrix (Im2) can be determined. Consider, for example, a karst 279 aquifer with a single conduit 0.2 m in diameter. The amount of water gathered by the conduit per unit length is assumed to be  $3 \times 10^{-7}$  m<sup>2</sup>/s (compare Figure 3, arrow 1) inducing a gradient 280 281  $I_{m1}$  in the matrix, approximated by the head difference (assuming the matrix head near the 282 conduit is related to the conduit perimeter  $h_1 = \pi d/2$ ) perpendicular to the conduit divided by 283 the associated length, which was set equal to 100 m. This gradient depends on the matrix hydraulic conductivity for a sufficient length scale  $K_{m1}$ , which is assumed to be 1 x 10<sup>-5</sup> m/s 284 (compare Figure 3, arrow 2a). The inflow gradient I<sub>m2</sub> was calculated for the hydraulic 285 conductivity in the vicinity of the conduit  $K_{m2}$ , which is assumed to be 1 x 10<sup>-9</sup> m/s (compare 286 Figure 3, arrow 2b). Since the inflow gradient  $I_{m2}$  of 4 x  $10^2$  is significantly increased 287 compared to the matrix gradient of  $1 \times 10^{-2}$ , it is expected that for this specific parameter set 288 289 under the assumed conditions, matrix heads significantly exceed the conduit heads in order to 290 obtain the necessary hydraulic gradient between the conduit and the matrix.



Figure 3: Flow rates  $Q_m/L$  and  $Q_i/L$  for a karst system with a single conduit (d = 0.2 m) for 293 294 several hydraulic conductivities. With a given amount of water that is gathered by the conduit 295 (Step 1), one can determine the equivalent hydraulic gradient in the matrix towards the 296 conduit. In this example the gradient is approximated by the head difference (assuming  $h_0 =$ 297  $\pi d/2$ ) perpendicular to the conduit divided by the associated length, which was set equal to 100 m (Step 2a:  $K_{m1} = 1 \times 10^{-5}$  m/s). The necessary inflow gradient  $I_{m2}$  for the conduit can be 298 determined according to the hydraulic conductivity in the matrix in the vicinity of the conduit 299  $K_{m2}$  (Step 2b:  $K_{m2} = 1 \times 10^{-9}$  m/s). As the inflow gradient  $I_{m2}$  clearly exceeds the matrix 300 301 gradient, matrix heads will clearly exceed conduit heads for this specific parameter set under 302 the assumed conditions.

### 303 **3. Numerical Computation of Turbulent Flow in Karst Aquifers**

# 304 3.1. Hybrid Models

To account for the three conduit-flow components (Figure 3), a hybrid model may be applied, which couples the discrete conduit network and the continuous matrix flow field by a head-dependent linear transfer term [Barenblatt et al., 1960; Bauer et al., 2003]

308

9 
$$Q_{ex} = \alpha A_{ex} K_{m2} (h_m - h_c) = f_{ex} L K_{m1} (h_m - h_c)$$
(13)

310

where  $Q_{ex}$  represents the exchange flow rate [L<sup>3</sup>/T],  $\alpha$  [1/L] is a factor that depends on the 311 312 conduit geometry and may be interpreted as inverse distance related to the head difference,  $A_{ex}$  [L<sup>2</sup>] is the exchange area between conduit and matrix,  $h_m$  is the hydraulic head in the 313 314 matrix [L], and h<sub>c</sub> represents the hydraulic head in the conduit [L]. Water exchange may also 315 be described as a function of the matrix hydraulic conductivity K<sub>m1</sub>, which is used for the 316 continuum model, and the conduit length L [L] by use of a dimensionless exchange factor  $f_{ex}$ . Accordingly, the exchange factor allows controlling the conduit infiltration capacity Q<sub>i</sub>/L. 317 318 This could be necessary as the matrix hydraulic conductivity used for the continuum may be 319 higher than that in the vicinity of the conduit. The reason for this behavior is the scale 320 dependency of hydraulic properties [Kiraly, 2002], compare section 2.3. Consequently, the 321 necessary gradient to capture Q<sub>m</sub>/L (flow towards the conduit) by Q<sub>i</sub>/L (conduit infiltration 322 capacity) is large. This scaling issue can be addressed by an adequately chosen value of fex. If 323 fex exceeds a certain threshold value, no additional flow resistance between conduit and matrix 324 exists and flow is limited by the matrix only. Bauer et al. [2003] observed this behavior while 325 investigating the influence of the exchange factor on karst aquifer genesis. Further 326 information regarding the numerical computation of laminar and turbulent flow in discrete

327 structures coupled to a continuum can be found in the literature [e.g. Liedl et al., 2003;
328 Shoemaker et al., 2008a].

329 3.2. Turbulent Single-Continuum Models

330 Standard groundwater continuum models ignore turbulent flow conditions and, 331 therefore, potentially represent pipe flow by a Darcian approach. An exception is CFPM2 for 332 MODFLOW-2005 [Shoemaker et al., 2008a], which considers turbulent flow effects within 333 continuum cells based on an approach introduced by Halford [2000]. MODFLOW-2005 334 [Harbaugh 2005] computes three-dimensional laminar groundwater flow in the matrix with 335 the following equation

336

337 
$$\frac{\partial}{\partial x} \left( K_{xx} \frac{\partial h_m}{\partial x} \right) + \frac{\partial}{\partial y} \left( K_{yy} \frac{\partial h_m}{\partial y} \right) + \frac{\partial}{\partial z} \left( K_{zz} \frac{\partial h_m}{\partial z} \right) \pm \psi = S_s \left( \frac{\partial h_m}{\partial t} \right)$$
(14)

338

339 where K<sub>xx</sub>, K<sub>yy</sub>, and K<sub>zz</sub> are the hydraulic conductivities [L/T] along the x, y, and z axes 340 respectively,  $\psi$  is the external volumetric flux per unit volume and represents sources and/or 341 sinks of water [1/T], S<sub>s</sub> is the specific storage [1/L], and t is time [T]. In the case of turbulent 342 flow conditions, CFPM2 modifies the horizontal hydraulic conductivity as a nonlinear 343 function in terms of head gradient and critical head gradient. Initially, CFPM2 was intended 344 to account for turbulent flow in aquifers with large pores where the turbulence occurs at low Reynolds number (1 to 100) [Kuniansky et al., 2008; Shoemaker et al., 2008a, b]. Reimann et 345 346 al. [2011] modified CFPM2 to consider turbulent flow with a user defined flow exponent m 347 (compare equation 3) such that the code is potentially able to account for conduit-type flow in 348 continuum cells – CTFc. Subsequently, the differences between the existing and modified 349 CFPM2 approaches are briefly introduced.

Karst conduits are represented in CFPM2 by continuum cells, which are indicated by a very high hydraulic conductivity and adequately small spatial discretization. To obtain values for the hydraulic conductivity, the Hagen-Poiseuille equation (equation 5) can be combined with the Darcy equation (equation 2) yielding the conduit hydraulic conductivity for laminar conditions K<sub>lam,c</sub> [L/T]

355

357

As stated in the previous section 2.2., the transition from laminar to turbulent flow conditions and vice versa is defined by a critical Reynolds number ( $Re_C$ ) converted to a critical head difference  $\Delta h_{crit}$  [Shoemaker et al., 2008a]

361

$$362 \qquad \Delta h_{crit} = \frac{\operatorname{Re}_{C} v \Delta L}{K_{lam} d} \tag{16}$$

363

364 where  $\Delta L$  represents the length over which the head difference is measured [L] and K<sub>lam</sub> is the 365 laminar hydraulic conductivity of the continuum [L/T].

The approach implemented in CFPM2 for calculating horizontal turbulent hydraulic conductivities is based on an approach for simulating turbulent flow in the vicinity of a wellbore [Halford, 2000]. A dimensionless adjustment factor  $F_{adj}$  is used to adjust the laminar hydraulic conductivity [Shoemaker et al., 2008a]

370

373 where  $K_{turb}$  [L/T] represents the turbulent hydraulic conductivity of the continuum. The 374 adjustment factor is

375

376 
$$F_{adj} = \sqrt{\frac{\text{Re}_C}{\text{Re}}} = \sqrt{\frac{K_{lam}\Delta h_{crit}}{K_{turb}\Delta h}}$$
(18)

377

378 when  $\Delta h > \Delta h_{crit}$  and with  $F_{adi} = 1$  when  $\Delta h < \Delta h_{crit}$ .  $K_{turb}$  is determined iteratively. This 379 approach was found to be well suited for describing turbulent flow observed in permeameter 380 tests at low Reynolds numbers [Kuniansky et al., 2008]. The resulting relationship between 381 hydraulic gradient and discharge, however, is found to be different from that predicted by the 382 Manning or Colebrook-White equations (equations 8 and 6, respectively). While the latter 383 predict a power law with an exponent m = 2.0 (compare Figure 1), steady-state flow 384 simulations with CFPM2 yield a power-law exponent m = 1.5 [Reimann et al., 2011]. 385 Consequently, the calculation of the adjustment factor was modified in CFPM2 such that the 386 power-law exponent can be specified by the user [Reimann et al., 2011]. Briefly described, 387 the discharge calculated by the laminar flow equation (Darcy's law) and that calculated by the 388 turbulent flow equation (power law with an exponent m = 2.0) are equal at an intersection 389 point defined by  $Re_{C2}$  (Figure 1). For that reason, the following equation can be derived for 390 the turbulent regime

391

$$q = \sqrt{\frac{I_{C2}}{I}} K_{lam} I \tag{19}$$

393

where I is the dimensionless hydraulic gradient and  $I_{C2}$  is the dimensionless hydraulic gradient where flow transitions from turbulent to laminar. Therefore, the adjustment factor resulting in CTFc with a flow exponent m = 2.0 is obtained as

$$F_{adj} = \sqrt{\frac{I_{C2}}{I}}$$
(20)

399

400 when  $\Delta h > \Delta h_{crit}$  and with  $F_{adj} = 1$  when  $\Delta h < \Delta h_{crit}$ . Accordingly, the friction coefficient of 401 common turbulent flow equations ( $k_c$  for the Colebrook-White equation or n for the Manning 402 equation) is represented by the critical Reynolds number  $Re_{C2}$ . Because CTFc is based on the 403 generalized power law with m = 2.0 (e.g., the Manning equation equivalent), the critical 404 Reynolds number Re<sub>C2</sub> was computed according to the Manning equation. Hence, Re<sub>C2</sub> is 405 linked with the Manning resistance coefficient n, for which appropriate values are available 406 within the literature (compare Table 1), by the following relationship, which is easily derived 407 from equations (1), (2), (8), and (15)

408

410

411 To account for hysteresis when transitioning from laminar to turbulent flow, the 412 critical Reynolds number  $Re_{C1}$ , is expressed as a multiple of  $Re_{C2}$ 

413

414 
$$\operatorname{Re}_{C1} = \varphi \operatorname{Re}_{C2} \tag{22}$$

415

416 where  $\varphi$  represents the multiplier, which ranges from 1 up to approximately 6, compare 417 Kanda and Yanagiya [2008].

The technical implementation of this approach in CFPM2 and its verification using a single-conduit setup are described by Reimann et al. [2011]. In the present work the model is applied to a coupled conduit-matrix flow system and compared to the hybrid model CFPM1. The impact of the linear head-dependent transfer term to control the conduit inflow capacity,
which is used in hybrid models (compare section 3.1), is considered in CTFc by the
Horizontal Flow Barrier Package (HFB) for MODFLOW [Hsieh and Freckleton, 1993].

424

# 4. Karst Spring Response to Recharge Events in Coupled Conduit-Matrix Systems

425 Karst systems consist of highly conductive conduits interacting with a matrix (Figure 426 2). These conduits have a complex configuration and connectivity. Application of the model 427 approaches described herein for a field situation that includes the complexity and uncertainty 428 of a real catchment would be beyond the scope of this study. Hence, a model study in a 429 realistic but relatively simple scenario was performed to (1) compare the SCM approaches 430 (existing CFPM2 [Shoemaker et al., 2008a], modified CFPM2 resulting in CTFc [Reimann et 431 al., 2011], and MODFLOW-2005 [Harbaugh, 2005]) as well as HM approaches (CFPM1 432 [Shoemaker et al., 2008a]); (2) examine the influence of turbulent flow on karst spring 433 responses; (3) investigate the influence of the matrix properties on karst spring responses; (4) 434 examine the impact of the water transfer coefficient that controls the interaction of conduit 435 and matrix, and (5) investigate the influence of matrix spatial discretization on karst spring 436 responses.

437 4.1. Study Setup

A single conduit with a length of 1,200 m is coupled to a matrix system to examine the matrix-conduit hydraulic interaction on short-term karst-spring responses after recharge events [Birk et al., 2006]. The width of both columns  $\Delta x$  and rows  $\Delta y$  is set to 50 m except for cells adjacent to the karst conduit centered in the synthetic catchment, where the width of the rows decreases successively to 0.2 m (50, 35, 20, 10, 5.5, 2.5, 1.5, 0.4, 0.2 m respectively). The matrix is represented as a confined / unconfined layer with a thickness of 1,000 m. The

diameter d of the karst conduit is 0.2 m with a pipe roughness k<sub>c</sub> of 0.01 m, which is 444 approximately equal to a Manning's n of 0.0183 s/m<sup>(1/3)</sup> and therefore adequate for a small 445 446 natural flow structure (compare equation 10 and Table 1). For the continuum models, the 447 karst conduit is represented by highly conductive cells with  $\Delta y = 0.200$  m and a thickness  $\Delta z$ 448 = 0.157 m, i.e. the flow cross-section is equal to that of the pipe. The equivalent laminar 449 hydraulic conductivity K<sub>lam,c</sub> of these cells can be determined using equation 15 yielding 450 9,374.8 m/s. The critical Reynolds number Re<sub>C2</sub>, which defines turbulent flow, can be 451 computed by using equation 21 yielding 892.9. The model setup is illustrated in Figure 4.



Figure 4: Model domain with steady-state hydraulic heads with a single conduit coupled to the matrix (computed by CTFc with  $K_{m1} = 1 \times 10^{-5}$  m/s and  $S_m = 0.01$ ). Grey colored arrow indicates a fixed head boundary while other boundaries are no-flow boundaries.

456 The discharge area is represented by a fixed hydraulic head of 1 m at the left hand side of the 457 conduit, i.e. flow is from right to left. The other boundaries of the model domain are 458 Neumann boundaries (no-flow). The entire catchment is supplied with constant diffuse recharge of 6.39 x  $10^{-9}$  m/s. Starting at t = 0 with the steady-state flow field, a direct recharge 459 460 pulse is injected at the right hand end of the conduit with a rate of 0.02 m<sup>3</sup>/s over a period of  $t_{event} = 7,200$  seconds. Recharge to the matrix remains unchanged. For an isolated conduit (i.e. 461 not coupled to the matrix), the event recharge Q<sub>event</sub>(t) is immediately passed through resulting 462 463 in a spring response  $Q_{\text{noexch}}(t) = Q_{\text{event}}(t) + Q_{\text{base}}(t)$  where  $Q_{\text{base}}(t)$  denotes the spring discharge 464 under steady-state flow conditions without direct recharge. The calculated spring discharge  $Q_{model}(t)$  was normalized relative to the pre-recharge event base-flow as  $Q_{norm}(t) = Q_{model}(t)$ 465 /Q<sub>base</sub>(t). 466

# 467 4.2. Benchmark Comparison of Model Approaches

468 The outlined scenario was used to compare different modeling approaches for a 469 coupled conduit-matrix system. The following numerical models are applied:

470 (1) MODFLOW-2005 (SCM, laminar flow in the continuum calculated by Darcy's471 equation).

(2) CFPM1 (HM, laminar flow in the continuum calculated by Darcy's equation, laminar /
turbulent flow in a discrete pipe network is simulated by the Hagen-Poiseuille and the
Colebrook-White equation). As mentioned above, this approach has been shown to
correctly reflect the dual-flow behavior of karst aquifers [Birk et al., 2005; Hill et al.,
2010]. Hence, this model approach provides a reference to which other simulations are
compared.

478 (3) CFPM2 (SCM, laminar flow determined using Darcy's equation, while turbulent flow
479 in the continuum is estimated by adjusting the hydraulic conductivity as described by
480 Shoemaker et al. [2008a]).

(4) Modified CFPM2 to result in CTFc (SCM, laminar flow calculated by Darcy's
equation, while turbulent flow in the continuum is provided by a power law with an
exponent of m = 2.0 [Reimann et al., 2011]).

The matrix conductivity and matrix storage were set to  $K_{m1} = 1 \times 10^{-5}$  m/s and  $S_m =$ 484 485 0.01. The fex threshold value, which creates additional flow resistance between matrix and 486 conduit, was approximated as 12.5. The inverse distance ( $\alpha$  in equation 13) used to compute 487 the f<sub>ex</sub> threshold value was assumed as 0.05 m, which is one fourth of the cell-width where 488 nodes are coupled to the continuum. Accordingly, the exchange factor  $f_{ex}$  of the hybrid model 489 was set to a very large value of 125, which is approximately ten times the threshold value, to 490 minimize conduit-matrix flow resistance, because there is no equivalent parameter in the SCM approach. 491

492 Figure 5 shows a very close match of the results computed with the CTFc with m = 2.0493 and the CFPM1 hybrid model based on the Colebrook-White equation. Hence, CTFc 494 reproduces the dynamic response of a karst spring as well as the hybrid model. In contrast, 495 both MODFLOW-2005 and the existing CFPM2 yield less damped spring responses, i.e. the 496 spring discharge is overestimated relative to the discharge computed by the HM CFPM1. If 497 only laminar flow is considered (MODFLOW-2005) or if turbulent flow is represented by 498 gradient-discharge relationship that obeys a power law with an exponent m = 1.5 (CFPM2), 499 the continuum model fails to reproduce the dynamic response of the spring as predicted by a 500 hybrid model that provides a straightforward representation of the coupled conduit-matrix 501 flow system. This statement is not affected by the slight oscillation in the CFPM2 that results 502 immediately after the increase in spring discharge.



504 **Figure 5:** Model comparison for the synthetic catchment.

505 Successful performance of the CTFc approach is premised on (1) continuum cells that were 506 discretized to represent the area of the karst conduit; and (2) an adequate critical Reynolds 507 number Re<sub>C2</sub> for the transition from turbulent to laminar flow conditions, which is consistent 508 with the intended flow resistance (e.g. according to Manning's equation). Whether an 509 adequate discretization (premise 1) can be achieved in practical applications and how far this 510 affects the performance of the turbulent continuum model need to be further examined by 511 site-related model applications. Identification of an adequate value of Re<sub>C2</sub> (premise 2), in 512 principle, requires knowledge about the geometric and hydraulic properties of the karst 513 conduits (compare equation 20). Re<sub>C2</sub> also may be determined by model calibration in 514 practical applications where information about conduit properties is lacking.

# 515 4.3. Parameter Study

The parameter study aims to investigate the influence of the matrix hydraulic properties (section 4.3.1), the conduit-matrix interaction (sections 4.3.2 and 4.3.4), as well as the influence of the matrix spatial discretization on karst spring responses (section 4.3.3). To quantify the effect induced by parameter variation on spring discharge, the spring response of the isolated conduit  $Q_{noexch}$  and that of the conduit-matrix system  $Q_{exch}$  are compared using the ratio of the respective water volumes discharged at the spring within the event duration  $t_{event} =$ 7.200 seconds

523

524 
$$\lambda_{Q} = \frac{\int_{t=0}^{t_{event}} \left[ \mathcal{Q}_{exch}(t) - \mathcal{Q}_{base}(t) \right] dt}{\int_{t=0}^{t_{event}} \left[ \mathcal{Q}_{noexch}(t) - \mathcal{Q}_{base}(t) \right] dt}$$
(23)

525

526 where  $\lambda_Q$  is termed the signal transmission factor [-]. A smaller signal transmission factor 527 indicates less water is discharged to the spring due to interaction with the matrix. 528 Consequently,  $\lambda_Q = 1.0$  means that the matrix does not affect the discharged water volume 529 within the event duration.

# 530 4.3.1. Influence of Matrix Parameters for Water Transfer between Matrix and Conduit

531 To investigate the influence of matrix hydraulic properties on spring discharge, matrix 532 hydraulic parameters are considered by 100 different realizations of randomly generated K<sub>m1</sub> and  $S_m$  values with  $K_{m1}$  ranging from 1 x 10<sup>-8</sup> to 1 x 10<sup>-4</sup> m/s and  $S_m$  ranging from 0.005 to 533 0.05 [-]. Matrix hydraulic conductivities are assessed according to Kiraly [2002] assuming a 534 535 controlling length scale for the continuum of approximately 100 m. As before, the exchange 536 factor f<sub>ex</sub> of the hybrid model was set to a value of 125 to allow unhampered water transfer 537 and, therefore, to allow comparability of model results. Simulated spring flow responses 538 obtained using the SCM approaches CTFc and MODFLOW-2005 as well as the HM 539 approach CFPM1 were compared.

540 CTFc reproduces normalized discharge simulated by the hybrid model CFPM1 541 reasonably well (Figure 6a-c). The magnitude and timing of both the rise and recession in 542 normalized discharge are similar for both CTFc and the hybrid model CFPM1. Comparable 543 performance between CTFc and CFPM1 was robust over the range in parameter realizations 544 for  $K_{m1}$  and  $S_m$ . The laminar flow model MODFLOW-2005 overestimates peak normalized 545 discharge by about 100% due to the lack of turbulent flow.

The signal transmission factor  $\lambda_Q$  (equation 23) was analyzed with respect to hydraulic properties of the matrix. For the scenarios presented here, spring response was insensitive to matrix storage. However, spring responses simulated with CTFc and CFPM1 were sensitive to the hydraulic conductivity of the matrix. Similar to the normalized discharge comparison, the sensitivity of the signal transmission factor  $\lambda_Q$  to matrix hydraulic conductivity computed by CTFc and CFPM1 agree reasonably well (Figure 6a-b, d-e) over the range in parameter

- 552 realizations for  $K_{m1}$  and  $S_m$ . The lower the matrix conductivity the less the influence of matrix
- 553 flow on conduit flow, which is indicated by an increasing signal transmission factor  $\lambda_Q$
- 554 (compare Figure 6 d-f). The laminar MODFLOW-2005 greatly underestimates the sensitivity
- 555 of  $\lambda_Q$  to  $K_{m1}$  due to reduced exchange flow between conduit and matrix.



556

**Figure 6:** Variation of matrix hydraulic parameters; the grey arrows indicate the direction of decreasing  $K_{m1} / S_m$  values; (a) – (c) Spring discharge normalized with respect to the pre-event base-flow, normalized spring discharge computed as  $Q_{norm}(t) = Q_{model}(t)/Q_{base}(t)$ , (d) – (f) signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) and comparison of hybrid model results (CFPM1) to the continuum models CTFc (e) and MODFLOW-2005 (f), (g) – (i) steady-state matrix heads along the cross-section A-A' (compare Figure 4), and (j) – (l) matrix head 32

- 564 change at the end of the recharge period at t = 7,200 s. Results were computed using CFPM1,
- 565 CTFc, and MODFLOW-2005, respectively.

Matrix hydraulic heads perpendicular to the conduit along the cross-section A-A' 567 568 (Figure 4) were examined for laminar and turbulent spring flow simulations (Figure 6 g-i). 569 Under steady-state conditions, flow in the conduit is turbulent with Re at the spring  $\sim 40,000$ . 570 Due to turbulence, the conduit head is elevated for CTFc and CFPM1 as compared to 571 MODFLOW-2005 as indicated by the associated matrix head near the conduit. Besides this, 572 steady-state matrix heads are comparable for all models (Figure 6g-i). The pulse recharge in 573 turbulent flow conduits increased Re at the spring to  $\sim 41,700$  to 127,000 (depending on the 574 matrix parameters) and causes comparably steeper hydraulic gradients (see also Figure 1) resulting in increased flow from the conduit to the matrix. Hydraulic heads in the surrounding 575 576 matrix reflect this behavior, whereas the response of matrix hydraulic heads in the laminar 577 model is less pronounced than in the turbulent models (Figure 6j-1).

### 578 4.3.2. Limiting Exchange Flow

579 Commonly, HMs couple matrix and conduit using a linear water exchange factor 580 (equation 13). A priori, this exchange factor is not considered by SCMs. Volumetric exchange 581 between conduit cells and the matrix, however, can be limited by the Horizontal Flow Barrier 582 (HFB) Package [Hsieh and Freckleton, 1993]. This could be important as previous 583 investigations by Peterson and Wicks [2005] suggest that only a limited portion of water is 584 transferred from the conduit to the matrix. Hence, the influence of hybrid model exchange 585 factors fex on spring flows and matrix heads is further examined using CFPM1, CTFc and 586 MODFLOW-2005. Experimental simulations were performed to determine if hybrid model 587 exchange factors can be duplicated with SCM's using the HFB Package. To this end, the 588 exchange factor fex was systematically varied for CFPM1 experimental simulations. Based on the basic model setup with  $K_{m1} = 1 \times 10^{-5}$  m/s and  $S_m = 0.01$ ,  $f_{ex}$  ranged from 100 to 0.001 in 589 590 100 equal steps. For the SCMs CTFc and MODFLOW-2005, the exchange factor was transformed into equivalent hydraulic characteristics of the HFB package. Comparison of the equation implemented in the HFB package (compare Hsieh and Freckleton [1993]) and equation 13, which is employed by CFPM1, reveals that the barrier hydraulic conductivity divided by the width of the barrier between two cells, HYDCHR, can be computed as

595

596 
$$HYDCHR = \frac{f_{ex}K_{m1}}{h_{sat}}$$
(24)

597

where  $h_{sat}$  is the average saturated thickness of the conduit cell and the adjacent cell, which is computed by CFPM1. The result from this calculation has to be divided by two as flow barriers were positioned on both sides of conduit-representing cells, whereas in CFPM1 the exchange occurs only within the cells coupled to conduits.

602 Normalized spring discharge for CFPM1 and CTFc with HFB show a very close 603 match over the range of parameter realizations for the exchange factor fex. Normalized 604 discharge for the laminar MODFLOW-2005 with HFB clearly differs from turbulent 605 approaches. The recharge pulse is more or less unaffected as the pulse is routed through the 606 conduit (Figure 7a-c). The signal transmission factor illustrates the damping behavior related 607 to exchange factors for turbulent flow models (Figure 7d-f). If the exchange factor is larger 608 than a threshold value, strong damping of spring discharges occurs. The damping gradually 609 diminishes with decreasing exchange factor. If the exchange factor is low, the conduit and the 610 matrix are only weakly coupled and exchange flow between conduit and matrix is low. 611 Therefore, the recharge pulse is directly transmitted to the spring (Figure 7a-f) and results for 612 laminar and turbulent models are similar, at least for the setting considered here.



**Figure 7**: Variation of the exchange factor  $f_{ex}$ ; the grey arrows indicate the direction of decreasing  $f_{ex}$ ; (a) – (c) Spring discharge normalized with respect to the pre-event base-flow, normalized spring discharge computed as  $Q_{norm}(t) = Q_{model}(t)/Q_{base}(t)$ . (d) – (f) signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) and comparison of hybrid model results (CFPM1) to the continuum models CTFc (e) and MODFLOW-2005 (f),, (g) – (i) steady-state matrix heads along the cross-section A-A' (compare Figure 4), and (j) – (1) matrix head change at the end 36

- 621 of the recharge period at t = 7,200 s. Results were computed using CTFc, CFPM1, and
- 622 MODFLOW-2005, respectively.

Steady-state matrix heads along the cross section A-A' illustrate the meaning of the 624 625 exchange factor (Figure 7g-i). Large exchange factors allow extensive drainage resulting in 626 comparatively low matrix heads. With this, the conduit acts similar to an internal fixed head 627 boundary condition. If the exchange factor falls below the threshold, the additional resistance alters the conduit behavior from a fixed head to a flux-dependent head (Cauchy) condition. 628 629 The smaller the exchange factor the higher the matrix heads. This behavior is similar for 630 laminar and turbulent models. However, both laminar and turbulent approaches differ with 631 respect to the matrix head change induced by the recharge pulse (Figure 7j-1). For laminar 632 flow models, matrix heads around the conduit seem to be unaffected by the recharge pulse 633 whether exchange factors are high or low. In contrast, the matrix head changes near the 634 conduit are considerable for turbulent models (CFPM1 and CTFc-HFB), especially when 635 exchange factors  $f_{ex}$  are large (Figure 7j-l).

## 636 4.3.3. Sensitivity of Exchange Flow to Matrix Spatial Discretization

637 Next, the influence of the matrix spatial discretization was briefly investigated, as this 638 potentially affects the interaction between turbulent flow conduits and matrix cells. The 639 matrix spatial discretization of rows  $\Delta y$  was increased for CFPM1 and CTFc up to around 50 640 m whereas the conduit-representing cell in CTFc remains at  $\Delta y = 0.2$  m in order to allow 641 comparable flow computation. The HFB flow barrier characteristic was parameterized equal to the model run with close-meshed matrix discretization. Additionally, for one CFPM1 642 643 model scenario, the continuum-cell interacting with the conduit-cell was set to  $\Delta y = 0.2$  m as 644 used for the scenario with close-meshed matrix discretization. As in the previous model runs, 645 the exchange factor  $f_{ex}$  was varied from 100 to 0.001 in 100 equal steps.

646 Normalized discharge for model runs with larger cells clearly differs from the 647 previously investigated standard scenario (Figure 8a-f). If the conduit interacts with large 648 continuum cells (left hand side of Figure 8), water may enter a large aquifer volume without 649 appropriate resistance behavior resulting in some artificial damping (Figure 8a). The signal 650 transmission factor supports this finding as the values are very low for large exchange factors 651 (Figure 8d). With a decreasing exchange factor, water flow between the conduit and the 652 matrix is reduced and, therefore, the significance of matrix discretization is diminished. If the 653 discrete conduit interacts with comparably small continuum cells (middle column of Figure 654 8), water transfer is not artificially damped by a large and slow reacting continuum cell. 655 However, due to head-averaging in the large surrounding cells, the hydraulic gradient from 656 the conduit to the matrix is artificially somewhat reduced resulting in slightly enhanced water 657 flow perpendicular to the conduit that results in slightly more damping of conduit flow 658 (compare Figures 8b,e,h,k and 7a,d,g,j). CTFc-HFB shows a similar behavior as CFPM1 with 659 small interacting cells because turbulent conduit flow interacts with the matrix continuum in a 660 similar manner.



661

**Figure 8:** Influence of spatial discretization of the continuum model; the grey arrows indicate the direction of decreasing  $f_{ex}$ ; (a) – (c) Mean normalized discharge; (d) – (f) signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) as well as the difference between hydraulic heads in the matrix and the conduit. (g) – (i) matrix heads perpendicular to the conduit for steady-state conditions prior to the recharge pulse; (j) – (l) matrix head variation due to the recharge pulse. Results were computed using CTFc with HFB and CFPM1.

4.3.4. Influence of Matrix Parameters on Spring Flow in Dependency of the Exchange Factor

Next, the influence of the exchange factor for several model settings with varying matrix parameter settings was investigated. To this end, the exchange factor was varied for three CFPM1 model runs as 12.5, 0.125, and 0.025, representing three different scalings of the hydraulic conductivity, as mentioned above (section 3.1). Each model run consists of 100 random realizations for matrix hydraulic conductivity  $K_{m1}$  and storage  $S_m$  as described in section 4.3.1.

Normalized discharge generally increases as the exchange factor  $f_{ex}$  decreases (Figure 9a-c). As expected, decreasing the exchange factor  $f_{ex}$  gradually decouples the conduit and matrix flow interaction. As the exchange factor  $f_{ex}$  approaches zero, normalized discharge for the CFPM1 hybrid model resembles normalized discharge simulated by the laminar flow model MODFLOW-2005 (Figure 9a-c and Figure 6a-c).

A decreasing exchange factor  $f_{ex}$  changes the sensitivity relation between  $\lambda_Q$  and  $K_{m1}$  (Figure 681 682 9d-f). The relatively large exchange factor f<sub>ex</sub> equal to 12.5 increases the exchange flow and 683 therefore the sensitivity of  $\lambda_Q$  to K<sub>m1</sub>. More specifically,  $\lambda_Q$  is relatively sensitive to the entire 684 range of  $K_{m1}$  tested in this study when the exchange factor  $f_{ex}$  equals 12.5 (Figure 9d). The signal transmission factor  $\lambda_Q$  is relatively insensitive to  $K_{m1}$  less than about 1.0 x  $10^{-6}$  m/s, 685 686 when the exchange factor  $f_{ex}$  equals 0.125 (Figure 9e). Likewise,  $\lambda_0$  is relatively insensitive to  $K_{m1}$  less than about 1.0 x 10<sup>-5</sup> m/s, when the exchange factor  $f_{ex}$  equals 0.025 (Figure 9f). This 687 688 behavior is consistent as smaller exchange factors result in diminished exchange flow. This 689 effect is increased by decreasing matrix conductivity K<sub>m1</sub> (compare equation 13) because the 690 exchange factor in this study depends on this parameter.



691

**Figure 9:** Influence of exchange factor  $f_{ex}$  on model results for a wide variety of matrix parameters computed with CFPM1; the grey arrows indicate the direction of decreasing  $K_{m1}$  /  $S_m$  values; (a) – (c) Mean normalized discharge; (d) – (f) signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) as well as the difference between hydraulic heads in the matrix and the conduit. (g) – (i) matrix heads perpendicular to the conduit for steady-state conditions prior to the recharge pulse; (j) – (l) matrix head variation due to the recharge pulse.

Matrix heads clearly vary over the whole range of matrix hydraulic parameters (Figure 9g-i) with a variability exceeding several orders of magnitude. This effect is emphasized by decreasing the exchange factor. Use of a small exchange factor together with low hydraulic conductivity results in implausible matrix heads indicating that the reasonable range of parameters is exceeded (Figure 9g-i).

704 The influence of turbulent conduit flow on the variability of matrix hydraulic heads 705 due to the variability of conduit flow diminishes if the exchange factor is further reduced in 706 the hybrid model (Figure 9j-1). Again, the smaller the exchange factor the more the conduit 707 and matrix are hydraulically decoupled (Figure 9g-1). Matrix hydraulic heads in the 708 investigated model setting, therefore, depend on the transfer coefficient as well as on the 709 matrix hydraulic parameters (compare Figures 7g-i and 9g-i). In practical model applications, 710 the exchange factor therefore may be found useful for adjusting the model to measured 711 responses of the spring or the hydraulic heads in the matrix. This, however, bears the risk that 712 the exchange factor is utilized to compensate for deficiencies in the conceptual model or 713 errors in other model parameters such as net recharge and matrix hydraulic conductivity.

## 714 **5.** Conclusions

715 A new approach (CTFc) to simulate laminar and turbulent flow using a single-716 continuum model is tested for consistency with a more complex and physics-based hybrid 717 model approach. CTFc simulates turbulent flow using a power law with a flow exponent m =718 2.0. Analyses of the underlying laminar and turbulent flow equations allowed adequate CTFc 719 parameterization for testing purposes. With this, dual-porosity flow components of karstic 720 aquifers can be simulated within a single continuum where karst conduits are represented by 721 highly-conductive model cells. CTFc successfully simulates spring discharge of a conduit 722 system embedded within a porous-medium matrix.

723 CTFc results agree well with those computed by the hybrid model CFPM1, which 724 explicitly accounts for turbulent conduit flow using a discrete pipe network. A comparison of turbulent single-continuum and hybrid flow models with the traditional laminar 725 726 MODFLOW-2005 reveals the impact of turbulent flow on spring flow induced by recharge 727 events particularly for comparably high conductive matrix settings with effective conduit-728 matrix interaction. Spring-flow dynamics are found to be strongly influenced by conduit 729 hydraulics. Accordingly, the shape of the spring hydrograph predicted by a model that 730 accounts for turbulent flow conditions differs from that obtained with a laminar model 731 approach. For karst systems with a highly conductive matrix, a conventional laminar 732 MODFLOW-2005 model greatly overestimates peak spring discharge and underestimates 733 hydraulic gradients within the conduit. This results in reduced conduit flow interaction with 734 the matrix. Spring flow responses to recharge events are poorly damped when compared with 735 simulated estimates that account for transitional and turbulent flow.

736 Hybrid models couple discrete pipe-like conduits to a continuum domain and use an 737 exchange factor to control the transfer of groundwater between the matrix and the conduit 738 system. For sufficiently high exchange factors, simulation results obtained with CTFc are 739 found to be as adequate as those simulated with the hybrid model CFPM1. By reducing the 740 value of the exchange factor, however, hybrid models are able to consider limited hydraulic 741 interaction between karst conduits and the matrix, which may correspond to a reduced conduit 742 infiltration capacity, e.g., due to the scale dependency of hydraulic properties. CTFc mimics 743 the limited hydraulic interaction by use of the HFB package for MODFLOW, which is proven 744 here by 2D parameter studies. Therefore, the numerically simpler CTFc approach is believed 745 to offer a reasonable alternative to the more demanding hybrid models in practical 746 applications addressing karst aquifers.

More generally, this study suggests that inferences of aquifer properties from spring hydrographs are potentially impaired by ignoring turbulent flow effects. Therefore, adequate representation of turbulent flow conditions in karst models may deserve equal or greater attention than focusing on the general pros and cons of continuum and hybrid models discussed in the literature. It is admitted that discharge hydrographs simulated in this paper were highly simplified. Hence, future work will have to substantiate the conclusions drawn herein using more complex models of real karst catchments.

### 754 Acknowledgements

This work was funded by the Deutsche Forschungsgemeinschaft (DFG, German Research Foundation) under grants no. LI 727/11-1 and SA 501/24-1 and the Austrian Science Fund (FWF) under grant no. L576-N21. The authors express their appreciation to Robert A. Renken, Andrew J. Long, and Eric D. Swain of the USGS and Rudolf Liedl from the TU Dresden for their internal review prior to submission to the journal as well as Stephen R. H. Worthington and two anonymous reviewers for their insightful comments that greatly helped to improve the manuscript.

Sadly, we have to report that our colleague and friend Dr. Christoph Rehrl died on 21 March

763 2010 at the age of 37. His dedication and person were an inspiration to us all.

#### 764 **Reference List**

- Barenblatt, G. I., I. P. Zheltov, and I. N. Kochina (1960), Basic concepts in the theory of
  seepage of homogeneous liquids in fissured rock, Journal of Applied Mathematics and
  Mechanics (PMM), 24(5), 1286–1303.
- Bauer, S., R. Liedl, and M. Sauter (2003), Modeling of karst aquifer genesis: Influence of
  exchange flow, Water Resour. Res., 39(10), 1285, doi: 10.1029/2003WR002218.
- Bear, J. (1988), Dynamics of fluids in porous media. Repr. Dover (Dover books on physicsand chemistry), New York.
- Birk, S., T. Geyer, R. Liedl, and M. Sauter (2005), Process-based interpretation of tracer tests
  in carbonate aquifers, Ground Water, 43(3), 381–388.
- Birk, S., R. Liedl, and M. Sauter (2006), Karst spring responses examined by process-based
  modeling, Ground Water, 44(6), 832–836.
- Covington, M. D., C. M. Wicks, and Saar M. O. (2009), A dimensionless number describing
  the effects of recharge and geometry on discharge from simple karstic aquifers, Water
  Resour. Res., 45, W11410, doi:10.1029/2009WR008004.
- Eisenlohr, L., L. Király, M. Bouzelboudjen, and Y. Rossier (1997), Numerical simulation as a
  tool for checking the interpretation of karst spring hydrographs, Journal of Hydrology,
  193, 306–315.
- Forchheimer, P. (1901), Wasserbewegung durch Boden, Zeitschr. d. V. deutsch. Ing., 45,
  1782–1788.
- Ford, D. and P. Williams (2007), Karst Hydrogeology and Geomorphology, John Wiley &
  Sons Ltd, Chichester, England.
- Halford, K. J. (2000), Simulation and interpretation of borehole flowmeter results under
  laminar and turbulent flow conditions, in Seventh International Symposium on
  Logging for Minerals and Geotechnical Applications, 157–168. Golden, Colorado.

- Harbaugh, A. W. (2005), MODFLOW-2005. The U.S. Geological Survey Modular GroundWater Model the Ground-Water Flow Process, U.S. Geological Survey Techniques
  and Methods 6-A, Reston, Virginia.
- Hill, M. E., M. T. Steward, and A. Martin (2010), Evaluation of the MODFLOW-2005
  Conduit Flow Process, Ground Water, 48(4), 549-559, doi: 10.1111/j.17456584.2009.00673.x.
- Hsieh, P. A. and J. R. Freckleton (1993), Documentation of a computer program to simulate
  horizontal-flow barriers using the U.S. Geological Survey modular three-dimensional
  finite-difference ground-water flow model, U.S. Geological Survey Open-File Report
  92-477, 32 p.
- Jain, S. C. (2001), Open-channel flow, John Wiley & Sons, New York.
- Kanda, H. and T. Yanagiya (2008), Hysteresis curve in reproduction of Reynolds' color-band
  experiments, J. Fluid. Eng., 130(5), 051202, doi: 10.1115/1.2903741.
- Kaufmann, G. (2009), Modelling karst geomorphology on different time scales,
  Geomorphology, 106, 62-77, doi: 10.1016/j.geomorph.2008.09.016.
- Király, L. (1984), Régularisation de l'Areuse (Jura Suisse) simulée par modèle mathématique,
  in Hydrogeology of karstic terraines, edited by A. Burger, and L. Dubertret, 94–99,
  Heise, Hannover.
- Király, L. (2002), Karstification and Groundwater Flow. In: Gabrovšek, F. (ed.), Proceedings
  of the Conference on Evolution of Karst: From Prekarst to Cessation, PostojnaLjubljana, 155–190.
- Kuniansky, E. L., K. J. Halford, and W. B. Shoemaker (2008), Permeameter data verify new
  turbulence process for MODFLOW, Ground Water, 46(5), 768–771.

- Liedl, R., M. Sauter, D. Hückinghaus, T. Clemens, and G. Teutsch (2003), Simulation of the
  development of karst aquifers using a coupled continuum pipe flow model, Water
  Resour. Res., 39, 1057, doi: 10.1029/2001WR001206.
- Lindgren, R. J., A. R. Dutton, S. D. Hovorka, S. R. H. Worthington, and S. Painter (2004),
  Conceptualization and simulation of the Edwards Aquifer, San Antonio Region,
  Texas, Scientific Investigations Report, Reston, Virginia.
- 818 Meyer, B. A., Kincaid, T. R., and T. H. Hazlett (2008), Modeling karstic controls on
- 819 watershed-scale groundwater flow in the Floridian Aquifer of North Florida, in: Yuhr,
- 820 L. B., E. C. Alexander, Jr., and B. F. Beck (eds), Sinkholes and the engineering and
- 821 environmental impacts of karst, Geotechnical Special Publication No. 183, American
  822 Society of Civil Engineers, Reston, VA, 351-361.
- Muskat, M. (1946), The flow of homogeneous fluids through porous media, McGraw-Hill,
  Ann Arbor, Michigan.
- Painter, S. L., A. Sun, and R. T. Green (2006), Enhanced characterization and representation
  of flow through karst aquifers. Awwa Research Foundation; IWA Publ., London.
- Peterson, E. W. and C. M. Wicks (2005), Fluid and Solute Transport from a Conduit to the
  Matrix in a Carbonate Aquifer System, Mathematical Geology, 37 (8), 851–867.
- Peterson, E. W. and C. M. Wicks (2006), Assessing the importance of conduit geometry and
  physical parameters in karst systems using the storm water management model
  (SWMM), Journal of Hydrology, 329(1-2), 294–305.
- Rehrl, C., S. Birk, and A. B. Klimchouk (2008), Conduit evolution in deep-seated settings:
  Conceptual and numerical models based on field observations, Water Resour. Res., 44,
  W11425, doi: 10.1029/2008WR006905.

835	Reimann, T, S. Birk, C. Rehrl, and W. B. Shoemaker (2011), Modifications to the Conduit
836	Flow Process Mode 2 for MODFLOW-2005, Ground Water, doi: 10.1111/Fj.1745-
837	6584.2011.00805.x.
838	Scanlon, B. R., R. E. Mace, M. E. Barrett, and B. Smith (2003), Can we simulate regional
839	groundwater flow in a karst system using equivalent porous media models? Case
840	study, Barton Springs Edwards aquifer, USA, Journal of Hydrology, 276(1-4), 137-
841	158.
842	Şen, Z. (1995), Applied Hydrogeology for scientists and engineers, CRC Press Inc, Boca
843	Raton, Florida.
844	Shoemaker, W. B., E. L. Kuniansky, S. Birk, S. Bauer, and E. D. Swain (2008a),
845	Documentation of a Conduit Flow Process (CFP) for MODFLOW-2005, U.S.
846	Geological Survey Techniques and Methods Book 6, Chapter A, Reston, Virginia.
847	Shoemaker, W. B., K. J. Cunningham, E. L. Kuniansky, and J. Dixon (2008b), Effects of
848	turbulence on hydraulic heads and parameter sensitivities in preferential groundwater
849	flow layers, Water Resour. Res., 44, W03501, doi: 10.1029/2007WR006601.
850	Teutsch, G., and M. Sauter (1998), Distributed parameter modeling approaches in karst-
851	hydrological investigations, Bulletin d'Hydrogéologie, 16, 99–109.
852	Worthington, S. R. H. (2009), Diagnostic hydrogeologic characteristics of a karst aquifer

- 853 (Kentucky, USA), Hydrogeology Journal (17), 1665–1678.
- Young, D. F., B. R. Munson, and T. H. Okiishi (2004), A brief introduction to fluid
  mechanics, 3rd ed., Wiley, Hoboken, NJ.

#### 856 **Figure captions**

Figure 1: Hydraulic gradient and corresponding discharge for different linear (laminar flow) and nonlinear (turbulent flow) equations. The example is computed for a conduit with d = 0.2m,  $k_c = 0.01$  m, and  $n = 1.83 \times 10^{-2}$  s/m<sup>(1/3)</sup>.

860

Figure 2: Cross section of a karst conduit embedded in the matrix. The flow system is controlled by matrix flow to karst conduits  $(Q_m)$ , conduit infiltration capacity  $(Q_i)$ , and conduit flow capacity  $(Q_c)$ .

864

865 Figure 3: Flow rates  $Q_m/L$  and  $Q_i/L$  for a karst system with a single conduit (d = 0.2 m) for 866 several hydraulic conductivities. With a given amount of water that is gathered by the conduit 867 (Step 1), one can determine the equivalent hydraulic gradient in the matrix towards the 868 conduit. In this example the gradient is approximated by the head difference (assuming  $h_0 =$  $\pi d/2$ ) perpendicular to the conduit divided by the associated length, which was set equal to 869 100 m (Step 2a:  $K_{m1} = 1 \times 10^{-5}$  m/s). The necessary inflow gradient  $I_{m2}$  for the conduit can be 870 871 determined according to the hydraulic conductivity in the matrix in the vicinity of the conduit  $K_{m2}$  (Step 2b:  $K_{m2} = 1 \times 10^{-9}$  m/s). As the inflow gradient  $I_{m2}$  clearly exceeds the matrix 872 gradient, matrix heads will clearly exceed conduit heads for this specific parameter set under 873 874 the assumed conditions.

875

Figure 4: Model domain with steady-state hydraulic heads with a single conduit coupled to the matrix (computed by CTFc with  $K_{m1} = 1 \times 10^{-5}$  m/s and  $S_m = 0.01$ ). Grey colored arrow indicates a fixed head boundary while other boundaries are no-flow boundaries.

879

880 **Figure 5:** Model comparison for the synthetic catchment

881 Figure 6: Variation of matrix hydraulic parameters; the grey arrow indicate the direction of 882 decreasing  $K_{m1}$  /  $S_m$  values; (a) – (c) Spring discharge normalized with respect to the pre-event 883 base-flow, normalized spring discharge computed as  $Q_{norm}(t) = Q_{model}(t)/Q_{base}(t)$ , (d) – (f) 884 signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) and comparison of hybrid model results 885 886 (CFPM1) to the continuum models CTFc (e) and MODFLOW-2005 (f), (g) – (i) steady-state matrix heads along the cross-section A-A' (compare Figure 4), and (j) - (l) matrix head 887 888 change at the end of the recharge period at t = 7,200 s. Results were computed using CFPM1, 889 CTFc, and MODFLOW-2005, respectively.

890

891 Figure 7: Variation of the exchange factor f<sub>ex</sub>; the grey arrow indicate the direction of 892 decreasing  $f_{ex}$ ; (a) – (c) Spring discharge normalized with respect to the pre-event base-flow, 893 normalized spring discharge computed as  $Q_{norm}(t) = Q_{model}(t)/Q_{base}(t)$ . (d) – (f) signal 894 transmission factor (ratio between water volumes discharged by conduit-matrix system and 895 isolated conduit within recharge period) and comparison of hybrid model results (CFPM1) to 896 the continuum models CTFc (e) and MODFLOW-2005 (f), (g) - (i) steady-state matrix heads 897 along the cross-section A-A' (compare Figure 4), and (j) - (l) matrix head change at the end 898 of the recharge period at t = 7,200 s. Results were computed using CTFc, CFPM1, and 899 MODFLOW-2005, respectively.

900

Figure 8: Influence of spatial discretization of the continuum model; the grey arrow indicate the direction of decreasing  $f_{ex}$ ; (a) – (c) Mean normalized discharge; (d) – (f) signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) as well as the difference between hydraulic heads in the matrix and the conduit. (g) – (i) matrix heads perpendicular to the conduit for steady-state 906 conditions prior to the recharge pulse; (j) - (l) matrix head variation due to the recharge pulse. 907 Results were computed using CTFc with HFB and CFPM1.

**Figure 9:** Influence of exchange factor  $f_{ex}$  on model results for a wide variety of matrix parameters computed with CFPM1; the grey arrow indicate the direction of decreasing  $K_{m1}$  /  $S_m$  values; (a) – (c) Mean normalized discharge; (d) – (f) signal transmission factor (ratio between water volumes discharged by conduit-matrix system and isolated conduit within recharge period) as well as the difference between hydraulic heads in the matrix and the conduit. (g) – (i) matrix heads perpendicular to the conduit for steady-state conditions prior to the recharge pulse; (j) – (l) matrix head variation due to the recharge pulse.

915

### 916 **Table captions**

917 **Table 1:** Values for Manning coefficient n.